

Formation of mixed-phase particles during the freezing of polar stratospheric ice clouds

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Polar stratospheric clouds (PSCs) are extremely efficient at catalysing the transformation of photostable chlorine reservoirs into photolabile species, which are actively involved in springtime ozone-depletion events. Why PSCs are such efficient catalysts, however, is not well understood. Here, we investigate the freezing behaviour of ternary HNO₃-H₂SO₄-H₂O droplets of micrometric size, which form type II PSC ice particles. We show that on freezing, a phase separation into pure ice and a residual solution coating occurs; this coating does not freeze but transforms into glass below ~150 K. We find that the coating, which is thicker around young ice crystals, can still be approximately 30 nm around older ice crystals of diameter about 10 μm. These results affect our understanding of PSC microphysics and chemistry and suggest that chlorine-activation reactions are better studied on supercooled HNO₃-H₂SO₄-H₂O solutions rather than on a pure ice surface.

PSCs have a crucial role in ozone depletion in the winter/spring polar stratosphere¹⁻³ because chlorine-activation reactions on cloud particles proceed faster than those in the gaseous phase^{4,5}. According to current knowledge, PSCs can be one of three types, Ia, Ib and II, and are thought to consist of solid nitric acid trihydrate (HNO₃·3H₂O (NAT)) particles, supercooled liquid HNO₃-H₂SO₄-H₂O droplets and solid ice particles, respectively⁵⁻⁷. Type Ia and Ib PSCs are formed in the Arctic and Antarctic stratospheres. Type II PSCs are formed mainly in the cold Antarctic region below the ice frost-point of ~189 K by the homogeneous freezing of HNO₃-H₂SO₄-H₂O droplets⁸ with an excess of HNO₃. Herein we deal with the formation and morphology of type II PSC ice particles. Previously it had been believed that the freezing of HNO₃-H₂SO₄-H₂O droplets led to the formation of completely solid ice particles with the acids buried within the ice phase. This belief could be because the freezing of HNO₃-H₂SO₄-H₂O droplets rich in HNO₃ of size and composition representative of the polar stratosphere takes place at temperatures much lower than the eutectic point of the HNO₃-H₂O system (~231 K). Therefore, the fate of H⁺, NO₃⁻ and SO₄²⁻ ions during the freezing of HNO₃-H₂SO₄-H₂O droplets with an excess of HNO₃ has not been investigated before.

It was shown recently that the rate of chlorine-activation reactions may depend on the phase state of the surface of PSC ice crystals⁹. The first stage of the most important chlorine-activation reaction of HCl + ClONO₂ → Cl₂ + HNO₃ (refs 2,4) is the incorporation of HCl into the ice surface with subsequent HCl ionization². For the ionization of HCl about six or more H₂O molecules are needed^{10,11}, and so the rate of the reaction clearly depends on whether the surface of the ice is solid or coated with a quasi-liquid layer. The formation of the quasi-liquid layer through the interaction of HCl with ice was predicted³ and also observed experimentally^{9,12-14}. The chlorine-activation reactions on the surface of hexagonal ice were studied intensively, both theoretically and experimentally^{6,7,9-13,15,17}. The studies were motivated by the belief, as mentioned above, that the freezing of HNO₃-H₂SO₄-H₂O droplets produces completely solid ice particles. In this Article,

we present differential scanning calorimetry (DSC) results that demonstrate that during the freezing of HNO₃-H₂SO₄-H₂O droplets of size and composition representative of the polar stratosphere, H⁺, NO₃⁻ and SO₄²⁻ ions are expelled from the ice lattice to form a residual freeze-concentrated solution that does not freeze, but transforms into glass at ~150 K (refs 18-20). Our results indicate that type II PSC ice crystals cannot be completely solid, as previously believed, but are enveloped by a supercooled HNO₃-H₂SO₄-H₂O coating.

Results

Phase behaviour of HNO₃-H₂SO₄-H₂O droplets on cooling and/or warming. Figure 1 demonstrates the phase separation into ice and a residual solution during the freezing of HNO₃-H₂SO₄-H₂O droplets. In Fig. 1, the calorimetric thermograms of emulsified 23 weight per cent (wt%) HNO₃ and 3 wt% H₂SO₄ are presented. In the thermograms, the exothermic peaks (pointing upwards) are a result of the heat of fusion released during freezing. Figure 1a demonstrates that ice is formed at ~189 K on cooling. This is inferred because the melting peak at ~248 K exactly matches the melting point of ice in the phase diagram of HNO₃-H₂SO₄-H₂O that contains 3 wt% H₂SO₄ (see Fig. 2b). In Fig. 1b, the cooling thermogram demonstrates that, after the ice has formed, the glass transition¹⁸⁻²⁰ occurs at $T_g \approx 150$ K. The appearance of the glass transition indicates that the droplets do not freeze completely. After freezing, a fraction of each droplet remains liquid until its transformation into glass. The liquid, which remains unfrozen, is a residual solution of the H⁺, NO₃⁻ and SO₄²⁻ ions expelled from the ice lattice during freezing. The residual solution, being a viscous, fragile liquid^{18,19}, transforms into glass at $T_g \approx 150$ K. That the glass transition occurs at T_g much below the lowest atmospheric temperature (~183-185 K) indicates that the freshly formed PSC ice crystals can be enveloped with a residual HNO₃-H₂SO₄-H₂O coating. Previously, we reported on the formation of H₂SO₄-H₂O and H₂SO₄-HNO₃-H₂O coatings around cirrus ice crystals formed by the homogeneous freezing of H₂SO₄-H₂O

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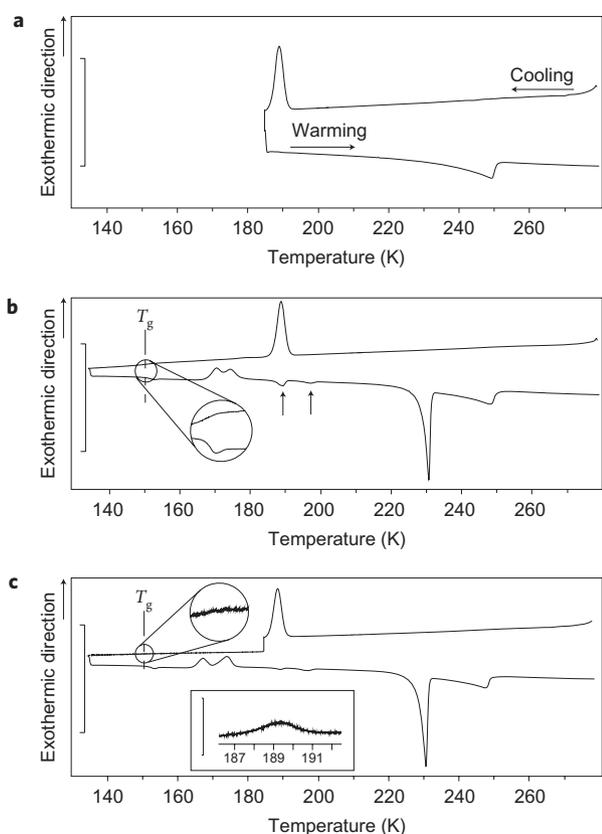


Figure 1 | DSC thermograms obtained from three samples of emulsified 23 wt% HNO_3 and 3 wt% H_2SO_4 . **a**, When cooling to and warming from ~ 185 K at 3 K min^{-1} the exothermic and endothermic peaks obtained result from the freezing and melting of ice, respectively. **b**, When cooling to ~ 133 K at 3 K min^{-1} additional transitions are seen, namely a liquid-to-glass transition^{18–20} on cooling and a glass-to-liquid transition on warming (magnified part). These transitions manifest themselves as a step in the corresponding thermogram^{20,23}. On warming, two additional freezing events and three additional melting events are observed (see text for explanation). Arrows point to the more subtle features of the glass-to-liquid transition and to two melting endotherms. **c**, The absence of a second freezing event on cooling and the presence of the liquid-to-glass transition (magnified part) is also found for a cooling rate of 0.05 K min^{-1} , which is more relevant to the polar stratosphere. The straight line between 185 and 133 K and the thermogram in the inset were obtained at 0.05 K min^{-1} . The scale bars (0.2 W g^{-1} in main plots; 0.01 W g^{-1} in inset) indicate heat flow through the samples.

droplets²¹ and H_2SO_4 – HNO_3 – H_2O droplets with an excess of H_2SO_4 (ref. 22). We also substantiated why the residual solution exists around ice crystals instead of in pockets within the ice²¹. In contrast to the H_2SO_4 – H_2O coating and the coating of H_2SO_4 – HNO_3 – H_2O with an excess of H_2SO_4 , which both freeze below ~ 183 K (refs 21–23), the HNO_3 – H_2SO_4 – H_2O coating with an excess of HNO_3 does not freeze, but undergoes a transition into glass at ~ 150 K.

Figure 1b shows that on warming, the glassy residual solution transforms at $T_g \approx 150$ K into a highly viscous liquid²³, which crystallizes at ~ 169 and ~ 173 K into ice and α -NAT, respectively. The α -NAT is thought to be metastable and converts into stable β -NAT at higher temperatures²⁴. The order of crystallization events was determined from additional DSC measurements, which showed that the peak at ~ 169 K indeed resulted from the crystallization of ice. Further warming produced two small melting peaks at ~ 189 and ~ 197 K (arrows in Fig. 1b) and two large peaks at ~ 231 and ~ 248 K. The existence of four melting events indicates

that four solids were formed on cooling and warming. As discussed above, ice surrounded by the freeze-concentrated solution melts at ~ 248 K, but our additional measurements reveal that other droplets with varying HNO_3 and constant H_2SO_4 concentrations also produce transitions similar to those presented in Fig. 1. The transition temperatures of these measurements are collected in Fig. 2. That the melting temperatures of ~ 189 , ~ 197 and ~ 231 K exist for droplets of different compositions suggests that these are three eutectic melting points. (1) The temperature of ~ 189 K may be the eutectic melting of ice– HNO_3 – H_2SO_4 – H_2O (ice–ternary hydrate). However, there are contradictory reports about the existence of an HNO_3 – H_2SO_4 – H_2O hydrate²⁵. (2) The eutectic mixture of ice–sulfuric acid tetrahydrate ($\text{H}_2\text{SO}_4 \cdot 4\text{H}_2\text{O}$ (SAT)) melts at ~ 197 K. In the H_2SO_4 – H_2O phase diagram^{26,27}, the eutectic temperature of ice–SAT is ~ 199 K. Our lower temperature results from the presence of HNO_3 . (3) The eutectic mixture of ice–NAT melts at ~ 231 K (ref. 25). The eutectic melting of ice–NAT decreases with increasing H_2SO_4 , as shown by the dashed green line in Fig. 2b. The dashed line is a locus of the eutectics of ice–NAT in bulk HNO_3 – H_2SO_4 – H_2O (ref. 25). Fig. 2b also shows that the concentration along the locus is ~ 32 wt% $\text{HNO}_3 + \text{H}_2\text{SO}_4$. After melting of ice–NAT, the only solid that remains is ice, so the concentration of residual solution is that along the locus. For the HNO_3 – H_2SO_4 – H_2O droplets with 0, 3 and 6 wt% H_2SO_4 studied here, the concentration of residual solution is $\sim 32/0$, $\sim 29/3$ and $\sim 26/6$ wt% $\text{HNO}_3/\text{H}_2\text{SO}_4$, respectively.

The cooling rate of 3 K min^{-1} (180 K h^{-1}) is larger than that encountered in the polar stratosphere (2 – 80 K h^{-1} (ref. 28)). This faster cooling rate may produce a freezing temperature of ice, T_i , colder than that in the polar stratosphere⁸. It may also cause a glass transition in the residual solution, rather than freezing. It is known that T_i decreases with an increasing cooling rate⁸ and the volume of a liquid determines a critical cooling rate, which ensures that freezing is avoided and brings about the glass transition²⁹. Therefore, we carried out DSC measurements using a cooling rate of 0.05 K min^{-1} (3 K h^{-1}), which is similar to the synoptic temperature change²⁸. In Fig. 1c, the inset demonstrates that the T_i obtained at 0.05 K min^{-1} is only ~ 1 K warmer than that obtained at 3 K min^{-1} . Obviously, the signal-to-noise ratio strongly decreased. The existence of the glass-to-liquid transition at $T_g \approx 150$ K in the warming thermogram in Fig. 1c implies that the residual solution undergoes the glass transition even if a cooling rate as small as 0.05 K min^{-1} is applied. To our knowledge, the finding that the glass transition occurs at such a small cooling rate is reported here for the first time. For stratospheric implications, the most important point is that the residual solution does not freeze at the same rate as the stratospheric cooling rate.

Coating of ice with a residual freeze-concentrated solution in the polar stratosphere. In Fig. 2a, the intersection of the yellow region with the experimental ice-freezing curves gives the concentration range of the droplets that would produce stable ice at polar stratospheric conditions, namely ~ 26 – 29 wt% HNO_3 for 0 wt% H_2SO_4 and ~ 23 – 25 wt% HNO_3 for 3 wt% H_2SO_4 . For simplicity, we calculated the thickness of the coating formed around ice crystals immediately after the freezing of 26–29 wt% HNO_3 droplets. The presence of a small amount of H_2SO_4 only slightly changed the thickness of the coating. We showed (see above) that after freezing HNO_3 – H_2O droplets, the residual solution was about ~ 32 wt% HNO_3 . We assumed that the diameter of HNO_3 – H_2O droplets is 0.4 – $2.4 \mu\text{m}$ and that freshly formed ice crystals are spherical with a smooth surface. In fact, as our optical microscopy observations show, the surface of ice is slightly rough (Fig. 3a). The assumption is justified because observation of cold cirrus clouds, which are formed near the tropopause region at similar conditions as PSCs, reveal quasi-spherical ice crystals of

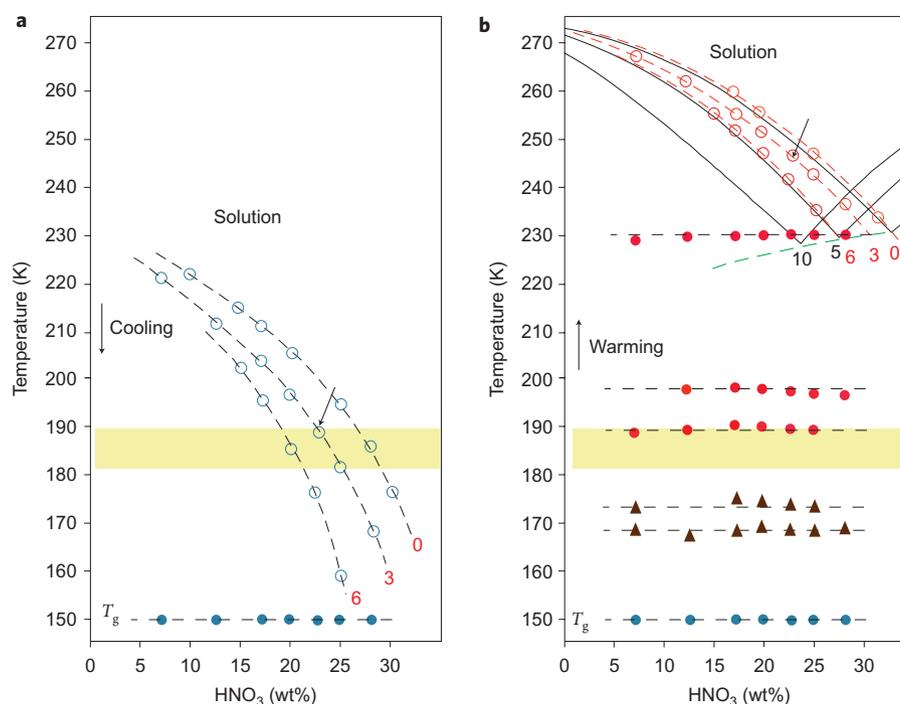


Figure 2 | The transition temperatures for the cooling and warming of bulk²⁵ and emulsified HNO₃-H₂SO₄-H₂O. **a**, The transition temperatures obtained on cooling. The red numbers denote wt% H₂SO₄ in the emulsified HNO₃-H₂SO₄-H₂O. The open and filled blue circles are the freezing and glass-transition points of emulsified HNO₃-H₂SO₄-H₂O. **b**, The transition temperatures obtained on warming. The red and black numbers denote wt% H₂SO₄ in the emulsified and bulk HNO₃-H₂SO₄-H₂O, respectively. The black solid lines are the phase diagrams of bulk HNO₃-H₂SO₄-H₂O (ref. 25). The green dashed line is a locus of the eutectics of bulk HNO₃-H₂SO₄-H₂O (ref. 25). The open red circles are the melting points of ice in emulsified HNO₃-H₂SO₄-H₂O. The solid red circles are the eutectics of ice-NAT (~231 K), ice-SAT (~197 K) and ice-ternary hydrate (~189 K) of the emulsified HNO₃-H₂SO₄-H₂O with 3 wt% H₂SO₄. The black triangles denote the cold crystallizations that occur on warming the emulsified HNO₃-H₂SO₄-H₂O with 3 wt% H₂SO₄ (see Fig. 1). In both **(a)** and **(b)** the arrows show the freezing and melting points of ice in the thermograms in Fig. 1. The yellow field is the temperature region of the stability of ice at the polar stratospheric conditions: 50 mbar pressure and a water mixing ratio of 5 parts per million by volume H₂O (ref. 5). Error bars are smaller than the size of the symbols.

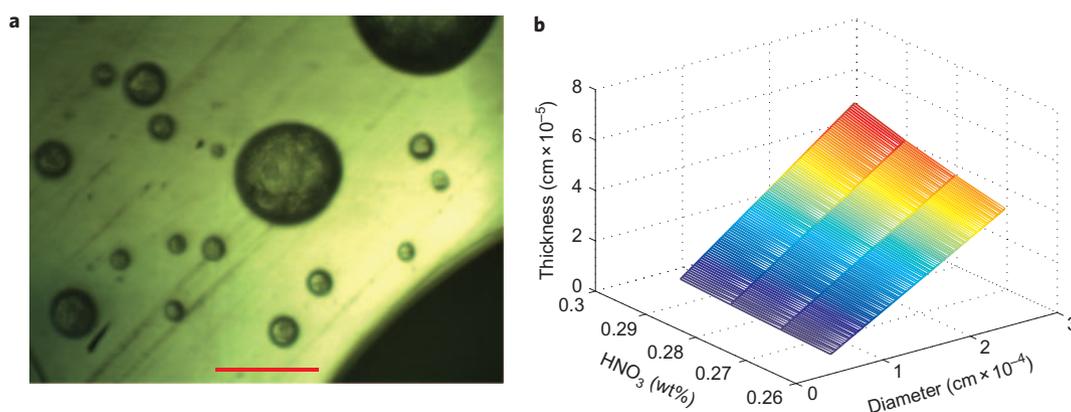


Figure 3 | Liquid coating around ice crystals. **a**, Microscopy image of mixed-phase particles (an ice core plus a solution coating) formed from 23/3 wt% HNO₃/H₂SO₄ droplets placed on a hydrophobic silicon surface. The picture was taken during warming at $T \sim 235$ K immediately after eutectic ice-NAT melting (see Fig. 1b). Scale bar = 50 μm . **b**, The initial thickness of the ~ 32 wt% HNO₃ coating that would envelop spherical ice crystals after freezing of 26-29 wt% HNO₃ droplets of diameter 0.4-2.4 μm .

diameter < 65 μm (ref. 30). We found that HNO₃-H₂O droplets produced an initial coating ~ 80 -600 nm thick (Fig. 3b).

Discussion

We have demonstrated that droplets of micrometric size experience a phase separation into pure ice and a residual freeze-concentrated solution at temperatures, chemical compositions and cooling rates relevant to the polar stratosphere. Freeze-concentrated,

HNO₃-rich ternary solutions do not freeze on continued cooling, but rather transform into a freeze-concentrated glass, in contrast to H₂SO₄-rich ternary solutions. As seen in the microscopy image in Fig. 3a, the liquid is expelled from the ice and coats it rather than being incorporated within the ice (for example, in pockets).

In the stratosphere, the surface of the outer coating equilibrates rapidly with the environment to reach the composition of freezing droplets. The inner surface of the coating remains in equilibrium

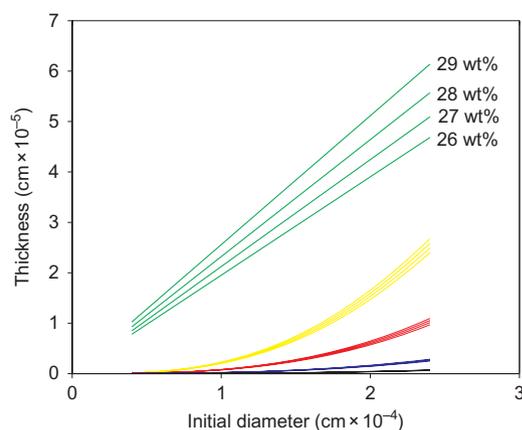


Figure 4 | Thickness of the ice crystal coating. The thickness of the coating that would envelop young and aged PSC ice crystals for four concentrations of HNO_3 in droplets from which the ice crystals formed. The green lines show the thicknesses of the initial coating (Fig. 3). The yellow, red, blue and black lines are the thicknesses of the coating around the aged ice crystals grown to diameters of 3, 5, 10 and 20 μm , respectively.

with the ice core, and thus preserves the composition of ~ 32 wt% HNO_3 . As a result of the concentration gradient, H_2O molecules on the surface diffuse through the coating and become incorporated into the ice. Assuming that the volume of the coating is limited by the acid content, the coating becomes thinner as the ice crystals grow. Ice PSC crystals can reach a diameter of ~ 10 μm (ref. 5). We calculated that the coating around such ice crystals would be ~ 30 nm thick (Fig. 4). In fact, the coating could be thicker because of the reactions of gaseous NO_y (ref. 6) and HCl (refs 11–13) with liquid H_2O . The reactions produce the condensed-phase HNO_3 (ref. 6) which, together with the additionally condensed H_2O , increases the volume of the coating and consequently its thickness. With time, as the amount of the condensed-phase HNO_3 increases, the coating may become almost binary HNO_3 – H_2O . Thus, our results indicate that chlorine-activation reactions on type II PSC ice crystals, which finally results in polar stratospheric ozone depletion in springtime, can be governed by supercooled HNO_3 – H_2SO_4 – H_2O or HNO_3 – H_2O solutions rather than by the pure ice surface. The coating may prolong the lifetime of ice crystals by suppressing their evaporation above ~ 189 K (ref. 24), which can account for the presence of ice crystals in the polar stratosphere at temperatures above 189 K (ref. 31).

Methods

We prepared HNO_3 – H_2SO_4 – H_2O solutions that contained 7–32 wt% HNO_3 and 0, 1, 3, 4 and 6 wt% H_2SO_4 by mixing 95–97 wt% H_2SO_4 and 65 wt% HNO_3 (Riedel-de Haën, Germany) with the corresponding amount of ultra-pure deionized water. Titration of HNO_3 – H_2O solutions against standard 2N NaOH showed an accuracy of ± 0.1 wt%. The oil phase used to form the emulsion was prepared by using 80 wt% Halocarbon 0.8 oil (Halocarbon Products Corp.) and 20 wt% lanolin (Sigma Aldrich). The oil phase and solution-in-oil emulsions were prepared according to the method of Chang *et al.* (ref. 8). The diameter of the droplets in the emulsions was measured with an optical microscope and found to be < 5 μm . In Fig. 2, only the experimental data for the droplets of composition 7–32 wt% HNO_3 and 0, 3 and 6 wt% H_2SO_4 are presented.

The calorimetric measurements of the emulsion samples were carried out with a Mettler Toledo DSC822. The calorimeter is well calibrated and reproduced the melting points of indium (429.75 K), water (273.15 K) and heptane (~ 182 K) with an accuracy of ± 0.4 K. The weight of the emulsion samples was ~ 25 – 35 mg. The DSC measurements were performed between 133 and 278 K at scanning rates of 3 and 0.05 K min^{-1} .

The oil-phase environment does not induce heterogeneous freezing of the droplets. This conclusion is inferred from a number of DSC measurements that showed emulsified, pure-water droplets always freeze at a homogeneous freezing temperature of $T_h \approx 233$ K. Chang *et al.* provide estimations of the stability of lanolin⁸, and showed no appreciable reaction between lanolin and concentrated

H_2SO_4 – H_2O and HNO_3 – H_2O on a time scale of two days⁸. Nevertheless, to prevent any degradation of lanolin, we made all DSC measurements in less than two hours from the moment of contact between the solutions and the oil phase. Repeated measurements performed for each composition of emulsified droplets showed very good reproducibility of the transition temperatures.

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Author contributions

A.B. designed the study, performed measurements and calculations, collected and analysed data, and wrote the manuscript. M.J.M. designed the study, analysed data, and wrote and corrected the manuscript. H.T., E.M. and T.L. analysed data, and wrote and corrected the manuscript. All authors discussed the results and commented on the manuscript.

Additional information

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