



Modeling the heterogeneous reaction probability for chlorine nitrate hydrolysis on ice

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[1] We present a theoretical estimate of the reaction probability γ for the chlorine nitrate (ClONO₂) hydrolysis on type II (water-ice) polar stratospheric cloud material. This estimate is based on high-level ab initio calculations in a supermolecule containing four molecules of water and one molecule of chlorine nitrate. To the best of our knowledge, this is the first estimate of γ that makes no a priori use of experimental data at all. Instead, the rate constants for association, surface reaction, and surface desorption as calculated by variational transition state theory enter the model. At 180 K we estimate $\gamma \approx 0.10^{+0.20}_{-0.06}$, which is within the error bars of the agreeing recommendations of Jet Propulsion Laboratory and International Union of Pure and Applied Chemistry of $0.3^{+0.7}_{-0.1}$. The temperature dependence between 75 and 150 K agrees with results obtained from laser-induced thermal desorption. In particular, the temperature of 105 K above which γ becomes less than unity is reproduced well. A negative temperature dependence between 180 and 210 K is found, which has not yet been confirmed in the laboratory for ClONO₂ hydrolysis but only for BrONO₂ on ice. This qualitative agreement of a gas-phase cluster calculation with experiments on hexagonal ice surfaces implies that a highly mobile and oxygen disordered ice surface rather than an ordered, immobile crystalline ice surface is experienced by chlorine nitrate molecules under polar stratospheric conditions.

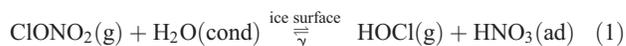
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1. Introduction

[2] Heterogeneous chemistry is recognized to be a key factor involved in polar ozone depletion for about 15 years [Solomon *et al.*, 1986; Molina *et al.*, 1987; Solomon, 1990]. Polar stratospheric clouds (PSCs) are known to catalyze the conversion of chlorine and nitrogen reservoir species (HCl, ClONO₂ and N₂O₅) efficiently (for a recent review see [Zondlo *et al.*, 2000]) and to be an effective sink of nitric acid (HNO₃). Type Ia PSCs are composed of nitric acid trihydrate (NAT), type Ib PSCs are supercooled ternary solutions of H₂SO₄/HNO₃/H₂O and type Ic PSCs are composed of small solid particles [Zhang *et al.*, 1993; Tabazadeh *et al.*, 1994; Toon *et al.*, 2000; Tolbert and Toon, 2001]. Because of the presence of acidic components in type I PSCs some experimental work has been devoted to studying the chemistry of chlorine nitrate with large ionic water clusters [Schindler *et al.*, 1996, 2000; Gilligan *et al.*,

2000; Gilligan and Castleman, 2001]. This study, however, focuses on type II PSCs, which are water-ice clouds occurring presumably as crystals of hexagonal ice.

[3] For chlorine reservoir species, three main activation reactions have been identified [Kolb *et al.*, 1995]. It has been shown that these three reactions likely occur efficiently also on the surface of cirrus clouds, thereby affecting the ozone depletion in the near-tropopause region [Solomon *et al.*, 1997]. Two out of these three are the conversion of HOCl or ClONO₂ by HCl on ice surfaces to molecular chlorine (Cl₂). This article, however, deals with equation (1), the hydrolysis of chlorine nitrate (ClONO₂) on water-ice surfaces:



[4] As early as 1986 it was recognized that the heterogeneous reaction on generally inert surfaces is more than 50 times faster than the homogeneous reaction [Rowland *et al.*, 1986]. In the following years the hydrolysis has been studied extensively using various experimental techniques, e.g., flow tubes coupled to mass spectrometers, and employing different surfaces, e.g., nitric acid hydrates [Moore *et al.*, 1990; Abbatt and Molina, 1992; Molina *et al.*, 1993; Zondlo *et al.*, 1998], sulfuric acid hydrates [Tolbert *et al.*,

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1988; Molina *et al.*, 1993; Zhang *et al.*, 1994], supercooled ternary solutions of H₂SO₄/HNO₃/H₂O [Molina *et al.*, 1993; Zhang *et al.*, 1995; Del Negro *et al.*, 1997], vapor-deposited ice [Leu *et al.*, 1991; Berland *et al.*, 1997], smooth ice [Lee *et al.*, 1999], and sodium chloride [see, e.g., Finlayson-Pitts *et al.*, 1989] or surfaces obtained from vapor codeposition on substrates [Koch *et al.*, 1997]. Concomitantly, also a lot of theoretical work has been done to resolve important details about the reaction mechanism [Ying and Zhao, 1997; Bianco and Hynes, 1998; Bianco *et al.*, 1998; McNamara and Hillier, 1999; McNamara *et al.*, 1999; Xu and Zhao, 1999; McNamara and Hillier, 2000; Bianco *et al.*, 2001; McNamara and Hillier, 2001; Loerting and Liedl, 2001; Bianco and Hynes, 2006]. These studies involved mainly molecular dynamics simulations and gas-phase cluster calculations by employing density functional theory or ab initio methods such as Møller-Plesset perturbation theory (MP2). Whereas the reactions are reasonably well understood from a phenomenological point of view due to the work in many research groups, there is still a vast amount of uncertainty in the microscopic view of the reaction mechanism. Especially the rates of the elementary steps in the overall process have not yet been characterized [Molina *et al.*, 1996]. However, a proper understanding of γ and the various underlying chemical and physical processes is required for a sound basis of interpreting and predicting ozone losses [Carslaw and Peter, 1997; Carslaw *et al.*, 1997].

[5] We therefore contribute to bridge this gap by calculating the reaction probability γ from first principles. Rate constants for the association, dissociation and reaction on the surface are determined from ab initio calculations and quantum transition state theory. This is, to the best of our knowledge, the first study that tries to model heterogeneous reaction probabilities without using experimental data at all. Experimental data is solely employed a posteriori for comparing with the theoretical prediction.

2. Methods

[6] The association rate constant $k_{\text{ass}}(T)$ (in Pa⁻¹s⁻¹) is calculated from [Tabazadeh and Turco, 1993; Adamson, 1997]

$$k_{\text{ass}}(T) = \frac{\alpha\sigma_0}{\sqrt{2\pi Mk_B T}} \exp\left(-\frac{\Delta G_{\text{ass}}^\ddagger(T)}{RT}\right) \quad (2)$$

where α is the mass accommodation coefficient of chlorine nitrate on ice (assumed to be one under stratospheric conditions), σ_0 is the area of one adsorption site on the surface (approximated to be 10⁻¹⁹ m²), M is the weight of a molecule of chlorine nitrate (1.618 × 10⁻²⁵ kg), k_B is the Boltzmann constant (1.381 × 10⁻²³ J K⁻¹), T the temperature (in K), R is the gas constant (8.315 J K⁻¹ mol⁻¹), and $\Delta G_{\text{ass}}^\ddagger(T)$ represents the temperature-dependent Gibbs free energy barrier to association (in J mol⁻¹), which is assumed to be zero as long as the Gibbs free energy barrier to desorption (dissociation) $\Delta G_{\text{des}}^\ddagger(T)$ has a positive sign. The latter is used to calculate the desorption rate constant $k_{\text{des}}(T)$ (in s⁻¹) [Pilling, 1995; Billing, 2000]:

$$k_{\text{des}}(T) = \frac{k_B T}{h} \exp\left(-\frac{\Delta G_{\text{des}}^\ddagger(T)}{RT}\right) \quad (3)$$

where h is the Planck constant (6.626 × 10⁻³⁴ J s). The desorption rate constant $k_{\text{des}}(T)$ is calculated from the simplified reaction scheme:



where n is the number of water molecules representing the ice surface. Ideally, one would like to use a hexagonal crystal for (H₂O)_{*n*}. However, as we are solving the time-independent Schrödinger equation using post-Hartree Fock methods and large basis sets in order to achieve the best possible description for the hydrogen bond interaction and dynamics, this task cannot be achieved because of limitations in computer time. These high-accuracy methods scale unfavorably with system size, so that incorporating $n = 4$ water molecules instead of $n = 3$ water molecules means an approximate 100fold increase of computer time for calculating a single-point energy (at 0 K) and a 1000-fold increase for calculating thermal corrections (including entropy contributions) to the single point energies. Thermal corrections are obtained at stationary points of the potential energy surface (PES), where the first-derivative with respect to nuclear coordinates (gradient) equals zero by calculating the Hessian matrix (second derivative). We estimate that a calculation at this level of theory using a state-of-the art high-speed computer would take several years for $n > 10$. We are using a chain of four water molecules representing the ice surface to which chlorine nitrate is bound via dangling OH bonds therefore and calculate the desorption barrier by removing chlorine nitrate from the surface at G2MP2 level of theory [Curtiss *et al.*, 1993]. Thermal corrections to the electronic energy are obtained at the hybrid density functional level of theory (B3LYP/6-31+G(d)) [Becke, 1993, 1997].

[7] The reaction rate constant $k_{\text{rxn}}(T)$ for the conversion of adsorbed chlorine nitrate to adsorbed HOCl and nitric acid is also calculated under the cluster assumption:



using high-accuracy quantum-mechanical methods. In particular G2(MP2) is employed to calculate the reaction barrier, and B3LYP/6-31+G(d) energies along the minimum energy path are interpolated to the G2(MP2) barrier. The details for the calculation on the full-dimensional PES (45 vibrational degrees of freedom for the $n = 4$ system) can be found elsewhere [Loerting and Liedl, 2001]. In brief, $k_{\text{rxn}}(T)$ (in s⁻¹) is calculated from quantum mechanical variational transition state theory [Chuang *et al.*, 1999; Corchado *et al.*, 1999; Frisch *et al.*, 1998]:

$$k_{\text{rxn}}(T) = \kappa(T) \frac{k_B T}{h} \exp\left(-\frac{\Delta G_{\text{rxn}}^\ddagger(T)}{RT}\right) \quad (6)$$

where $\kappa(T)$ is the enhancement factor of the reaction due to quantum mechanical tunnelling along all possible directions of the PES (of course dominated by contributions corresponding to proton motion) and $\Delta G_{\text{rxn}}^\ddagger(T)$ is the free energy barrier approximated by the zero-point corrected energy barrier along the minimum energy path (MEP) on the PES [Truhlar and Garrett, 1984; Truhlar *et al.*, 1985;

Truhlar and Gordon, 1990; Truhlar, 1995]. The rate for the backward reaction is also calculated with the same approach, and is found to be slower by several orders of magnitude at any given temperature. As in all common derivations, the influence of the back reaction on the uptake probability γ is neglected, therefore.

[8] Using such a data set of $k_{rxn}(T)$, $k_{des}(T)$, and $k_{ass}(T)$, i.e., without the use of experimental data, γ can be estimated from [Tabazadeh and Turco, 1993]

$$\gamma(T, p^{\text{ClONO}_2}) = \frac{4k_B T K k_{rxn}}{\sigma_0 \bar{c} (1 + K p^{\text{ClONO}_2})} \quad (7)$$

where \bar{c} is the average thermal velocity,

$$\bar{c} = \sqrt{\frac{8k_B T}{\pi M}} \quad (8)$$

p^{ClONO_2} is the partial pressure of chlorine nitrate (in Pa) and K (in Pa^{-1}) is calculated from rate constants only:

$$K = \frac{k_{ass}}{k_{rxn} + k_{des}} \quad (9)$$

K can be employed to calculate the surface coverage θ of ice according to the (surface reactive) Langmuir adsorption isotherm [Langmuir, 1922]:

$$\theta(p^{\text{ClONO}_2}) = \frac{K p^{\text{ClONO}_2}}{1 + K p^{\text{ClONO}_2}} \quad (10)$$

We base this model on the simple Langmuir isotherm on a similar basis as a previous model simulating the effect of competitive adsorption on PSC material [Mozurkewich, 1993]. Note that for $k_{rxn} \gg k_{des}$, equation (7) simplifies to

$$\gamma = \frac{4k_B T k_{ass}}{\sigma_0 \bar{c}} = \alpha \exp\left(-\frac{\Delta G_{ass}^\ddagger(T)}{RT}\right) \quad (11)$$

and is no longer temperature-dependent (as long as there is no association barrier and α is constant with temperature), and equals to 1 under the assumptions made for the stratosphere. On the other hand, for $k_{des} \gg k_{rxn}$, equation (7) simplifies to

$$\gamma(T) = \frac{4k_B T k_{ass} k_{rxn}}{\sigma_0 \bar{c} k_{des}} \quad (12)$$

and the temperature dependence of γ originates from the temperature dependence of $\frac{k_{rxn}}{k_{des}}$. Note that equation (7) can alternatively be written as

$$\gamma(T) = \alpha \frac{k_{rxn}}{k_{rxn} + k_{des}} \exp\left(-\frac{\Delta G_{ass}^\ddagger(T)}{RT}\right) \quad (13)$$

which reduces to

$$\gamma(T) = \frac{k_{rxn}}{k_{rxn} + k_{des}} \quad (14)$$

under the assumptions of $\alpha = 1$ and no association barrier.

[9] We set p^{ClONO_2} to 10^{-8} Pa, which is very close to the polar stratospheric partial pressure, and use equation (7) to

estimate γ and equation (10) to estimate θ . We note that a wide range of partial pressures of chlorine nitrate has been employed in different experimental studies. However, the estimate of γ does not depend much on the partial pressure of chlorine nitrate. In fact, under the assumptions of equation (14), γ does not depend on the partial pressure of chlorine nitrate. Therefore our data can be compared with experimental data, even though some flow tube studies employed higher partial pressures.

[10] The experimental results are influenced from differences in ice surface area, porosity and tortuosity [Keyser et al., 1991; Grassian, 2001] and surface poisoning by HNO_3 [Hanson and Ravishankara, 1991], which decreases γ by approximately a factor of 50 when the surface saturates with the product HNO_3 . Surface porosity might play a role in the magnitude of the uptake probability [Chu et al., 1993; Hanson and Ravishankara, 1993]. The recommendations given by International Union of Pure and Applied Chemistry/Jet Propulsion Laboratory (IUPAC/JPL) are based on a critical evaluation of the available experimental studies. In particular they have evaluated, to which degree the experimental studies suffer from HNO_3 contamination. Therefore the IUPAC/JPL recommendations provide the benchmark for our theoretical estimates of γ on a fresh ice surface.

3. Results and Discussion

3.1. Reaction Mechanism

[11] Figure 1 shows the chemical reaction mechanism, denoted by $n = 4$, that was found on minimizing the electronic energy as a function of the positions of the nuclei for a system containing four water molecules and one molecule of chlorine nitrate (as implemented by Frisch et al. [1998]). In comparison to the molecular systems studied previously by two of us (see Figure 1, $n = 1$ to $n = 3$) [Loerting and Liedl, 2001], there is one crucial difference: The reaction barrier is considerably lower, and amounts to $5.61 \text{ kcal mol}^{-1}$ (G2MP2 level of theory), which is a reduction of the barrier by $14.34 \text{ kcal mol}^{-1}$ as compared to the system with just one molecule of water less ($n = 3$). This reaction barrier agrees well with the best estimate given in the literature of $3\text{--}7 \text{ kcal mol}^{-1}$ [Tabazadeh et al., 1994]. Also the fact that on ice films kept between 90 and 120 K no reaction product was observed is indicative of a barrier for the surface reaction [Donsig et al., 1999]. We believe that the drastic reduction of the reaction barrier for the surface reaction is caused by the change of the angle of water molecules with respect to each other. Whereas for $n = 3$ the O-O-O angle amounts to 102.5° , it amounts to 120.0° for any combination of three H_2O oxygen atoms in the $n = 4$ model system for the ice surface- ClONO_2 interaction. All of these angles are reasonably close to the ideal tetrahedral angle of 109.5° observed in X-ray studies of hexagonal ice [Hobbs, 1974]. However, in the reaction product, and according to the Hammond postulate [Hammond, 1955] also to a certain degree in the transition state, the angle between three oxygen atoms adopts unfavorable values far away from 109.5° for the $n = 3$ model, but still reasonably close values for the $n = 4$ model. This situation allows shuttling of hydrogen atoms between oxygen atoms in a favorable near-linear configuration. Whereas in a true hexagonal crystal the angular flexibility would be significantly reduced, the cluster employed here shows flexibility in the optimized structures.

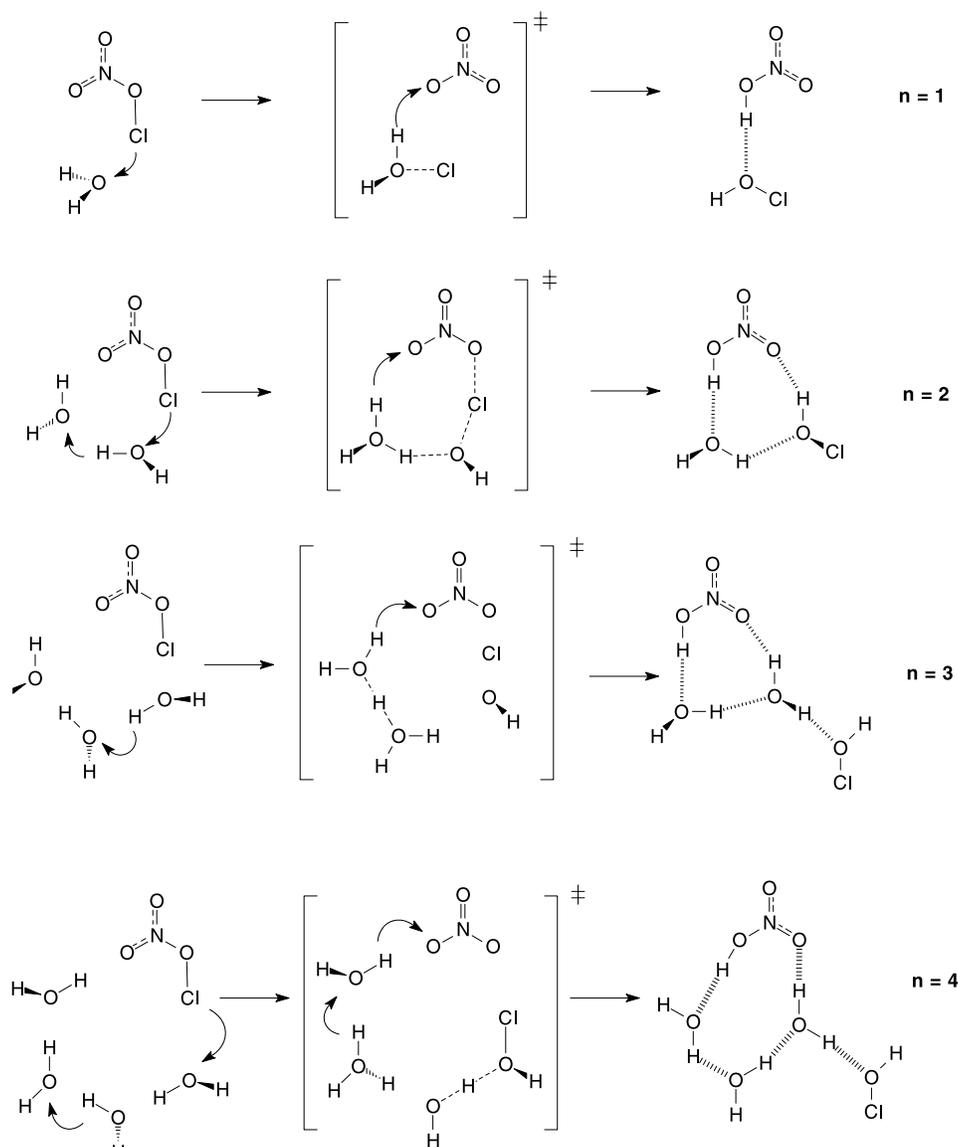


Figure 1. Chemical reaction mechanism for heterogeneous chlorine nitrate hydrolysis in clusters containing $n = 1$ to $n = 4$ water molecules, including ClONO_2 -water surface adsorbate, transition state to hydrolysis, and hydrogen-bonded reaction product of $\text{HOCl} \cdot \text{HNO}_3$ adsorbed to the surface. Note that for $n = 4$ the ice surface is represented by a water tetramer for the educts and reduced to a water trimer for the hydrolysis products. The angles between the water molecules resemble very much the crystallographic angles between water molecules in hexagonal ice, and the gas-phase cluster simplification of the reaction system results in the molecular flexibility that is presumably encountered at polar stratospheric conditions.

The combination of this flexibility and the near-hexagonal structure, to our opinion, results in the low-barrier situation that is also found in the stratosphere, and represented by both a JPL as well as a IUPAC-recommended estimate for the reaction probability at 200 K of $0.3^{+0.7}_{-0.1}$ [Hanson and Ravishankara, 1991; DeMore et al., 1997; Sander et al., 2000]. Of course, also microsolvation effects, which are not studied here, can contribute significantly to this flexibility [Bianco and Hynes, 1998]. It may seem awkward to think of oxygen atoms at the crystal surface as flexible. However, on type II polar stratospheric water-ice clouds such a flexibility could possibly be caused by a thin quasi-liquid layer, which may be induced by atmospheric species present in rather low

mixing ratios at temperatures of 185–210 K [Molina, 1994; McNeill et al., 2006].

3.2. Surface Reaction and Desorption

[12] In Figure 2 the rate constants for the hydrolysis $k_{\text{rxn}}(T)$ on the model surface and the rate constant for chlorine nitrate desorption $k_{\text{des}}(T)$ from the model surface employed for estimating the reaction probability γ are depicted. The $k_{\text{rxn}}(T)$ has been calculated from variational transition state theory as discussed in section 2. The $k_{\text{des}}(T)$ has been calculated by employing equation (3) and the free energy barriers for desorption as depicted in Figure 3.

[13] Because of a tunnelling enhancement of $k_{\text{rxn}}(T)$ by a factor of 2–4 at temperatures from 250 to 100 K (see Figure 4)

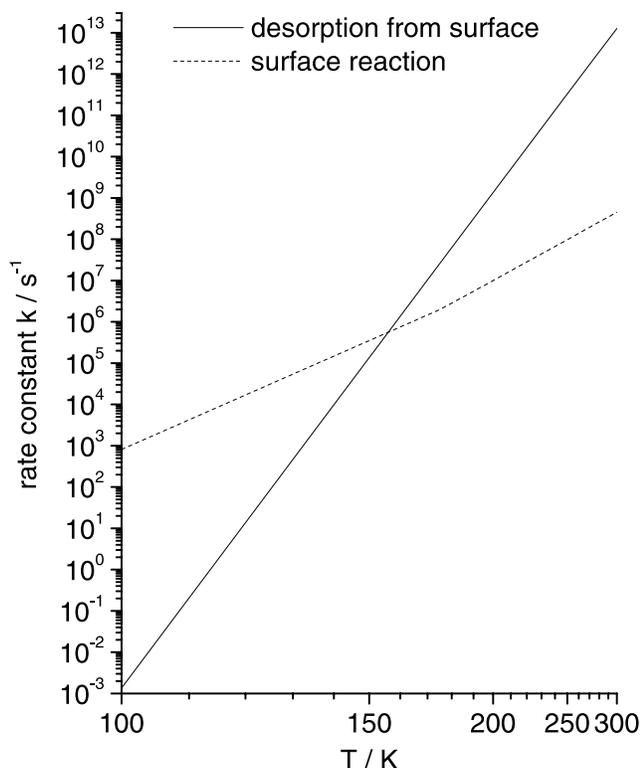


Figure 2. Rate constants of desorption from the model ice surface as compared to the rate constants of reaction on the model ice surface. Reaction rate constants k for the decomposition of chlorine nitrate in an (unimolecular) supermolecule containing four water molecules are calculated at a B3LYP/6-31+G(d) potential energy surface and interpolated to G2MP2 barriers as well as corrected for quantum effects. Rate constants for desorption were calculated according to equation (3) by employing the free energy barrier depicted in Figure 3.

the Arrhenius plot in Figure 2 is nonlinear. As compared to the model systems employing two or three water molecules, which show tunnelling enhancement factors reaching several orders of magnitude at 100 K (see Figure 4), the contribution of tunnelling is much smaller in the case of four water molecules. This can be attributed to enhanced movement of O atoms, which results in a higher reduced mass for proton transfer. At stratospheric temperatures the tunnelling effect is weak and comparable to the effect calculated for the homogeneous gas-phase reaction ($n = 1$) [Atkinson *et al.*, 1986]. So-called direct corner cutting, which causes a strong tunnelling effect, is important only at $T < 150$ K. It therefore does not play a significant role under tropospheric or stratospheric conditions.

[14] The rate constants calculated here compare to the estimate of 10^4 s^{-1} (200 K) to 10^6 s^{-1} (300 K) [Robinson *et al.*, 1997; Shi *et al.*, 2001]. This literature estimate is based on an extrapolation to pure water on the basis of experimentally measured values on sulfuric acid solutions. It is surprising that gas-phase clusters and an experimental extrapolation to pure water shows only a difference of about two orders of magnitude. We therefore conjecture that the hydrolysis in the gas-phase cluster and the hydrolysis in liquid water take place via a similar molecular mechanism.

[15] The rate constant of surface desorption $k_{\text{des}}(T)$ of chlorine nitrate on the other hand follows a near-linear Arrhenius behavior in the temperature region between 100 and 300 K, although the Gibbs free energy barrier to desorption $\Delta G_{\text{des}}^\ddagger(T)$ (compare Figure 3) significantly drops on increasing the temperature due to an increasing contribution of the translational entropy. Note that the barrier is neither sensitive to the cluster size nor sensitive to the quantum mechanical method employed, as judged from the near equality of the two approaches to the desorption barrier of chlorine nitrate.

[16] At all temperatures the barrier to desorption for HOCl is lowest, and becomes even negative at $T > 170$ K, which implies immediate desorption from the surface at, e.g., 200 K. The fact that HOCl is released to the gas phase on a timescale of minutes [Molina *et al.*, 1987] in some experiments may be related to the fact that before desorbing from the ice surface the hydrogen bond between HOCl and HNO_3 adsorbed on the surface has to be broken first. This retarding function of HNO_3 has also been mentioned lately from second harmonic generation studies on the basal ice surface of hexagonal ice [Geiger *et al.*, 2001], which showed also that HOCl desorbs immediately as long as the ice surface is relatively fresh and free of HNO_3 . Nitric acid by itself sticks to the surface at stratospheric temperatures, as shown in all experiments as well as by our calculations [Hanson and Ravishankara, 1992]. This sticky behavior is represented by the very high $\Delta G_{\text{des}}^\ddagger(\text{HNO}_3)$,

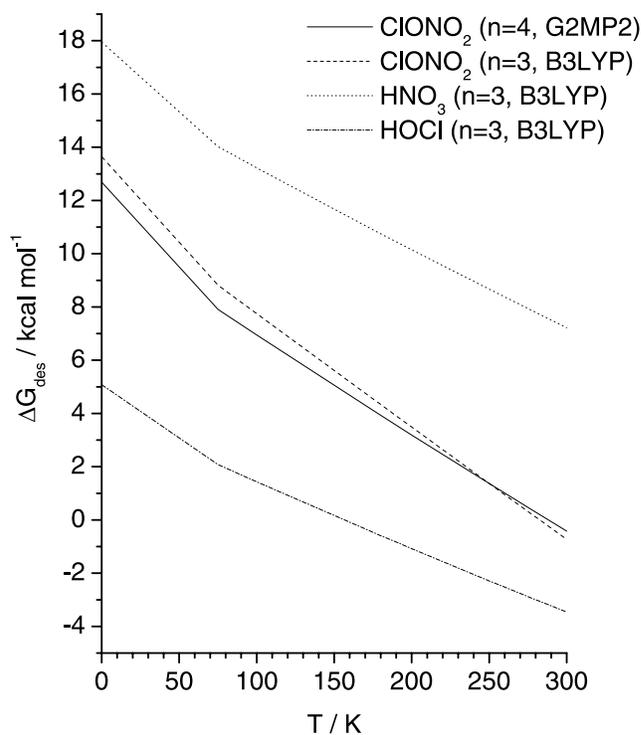


Figure 3. Gibbs free energy barriers of desorption from a model ice surface represented by either a linear water trimer or water tetramer as calculated at either G2MP2 or B3LYP/6-31+G(d) level of theory and thermally corrected by calculating second derivatives of the electronic energy with respect to nuclear coordinates, i.e., the Hessian matrix, at B3LYP/6-31+G(d) level of theory.

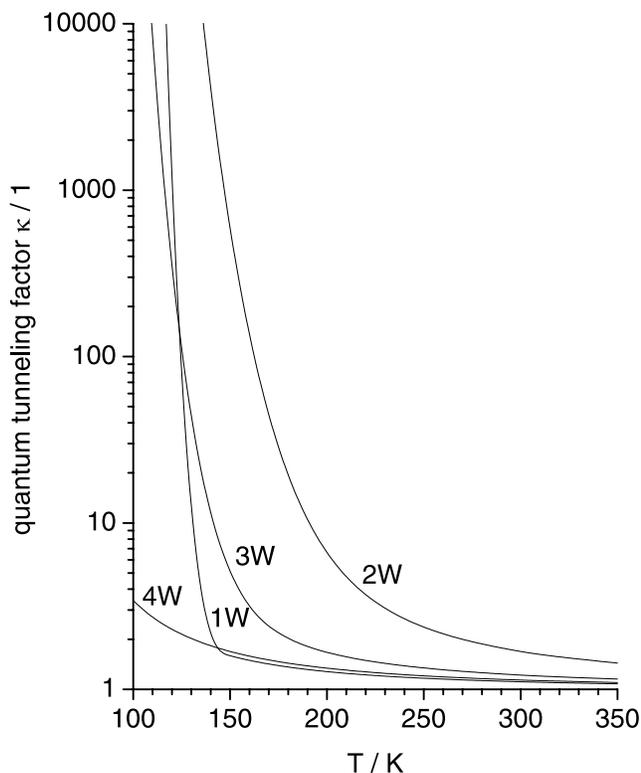


Figure 4. Tunnelling correction factors κ for the decomposition of chlorine nitrate in an unimolecular complex containing one (1W), two (2W), three (3W), or four (4W) water molecules as calculated at a B3LYP/6-31+G(d) potential energy surface under the microcanonical optimized multidimensional tunnelling approximation.

leading to desorption on a timescale much longer than hours. The desorption barrier of chlorine nitrate from the ice surface is intermediate between HNO_3 and HOCl . Similarly, *Bianco and Hynes* [2006] have found an activation barrier for desorption of ClONO_2 of 7 kcal mol^{-1} .

3.3. Ab Initio Estimate of the Reaction Probability

[17] As a result of the strong temperature dependence of the desorption rate constant and the much weaker dependence of the surface reaction rate constant these two curves cross each other at approximately 150 K. According to equation (14) this implies a reaction probability of 0.5 at the intersection point, which can also be seen in Figure 5. In Figure 5 the reaction probability γ (left ordinate, obtained from equation (7)) is shown together with the surface coverage θ (right ordinate, obtained from equation (10), $p^{\text{ClONO}_2} = 10^{-8} \text{ Pa}$). The surface coverage is far below monolayer coverage for all investigated temperatures, which justifies the use of the Langmuir isotherm and also allows us to use equation (13) instead of equation (7). These low coverages are in qualitative agreement with findings by *Hanson and Ravishankara* [1992], who detect a surface coverage clearly lower than 10^{-3} at 191 K.

[18] The estimate of γ drops from approximately 0.1 to 0.01 in the temperature range 180–210 K. At temperatures lower than 100 K γ is estimated to be unity, whereas it becomes lower than 0.0001 near the melting point of ice. A similar negative temperature dependence has been found in recent experiments, in which a temperature increase from ca. 120 K to 140 K was accompanied by a drop in γ of a factor of 6 [*Berland et al.*, 1997]. They have also found that below 105 K γ is independent of temperature and reaches the maximum value provided by the mass accommodation coefficient α of chlorine nitrate on an ice surface. Their curve displayed in

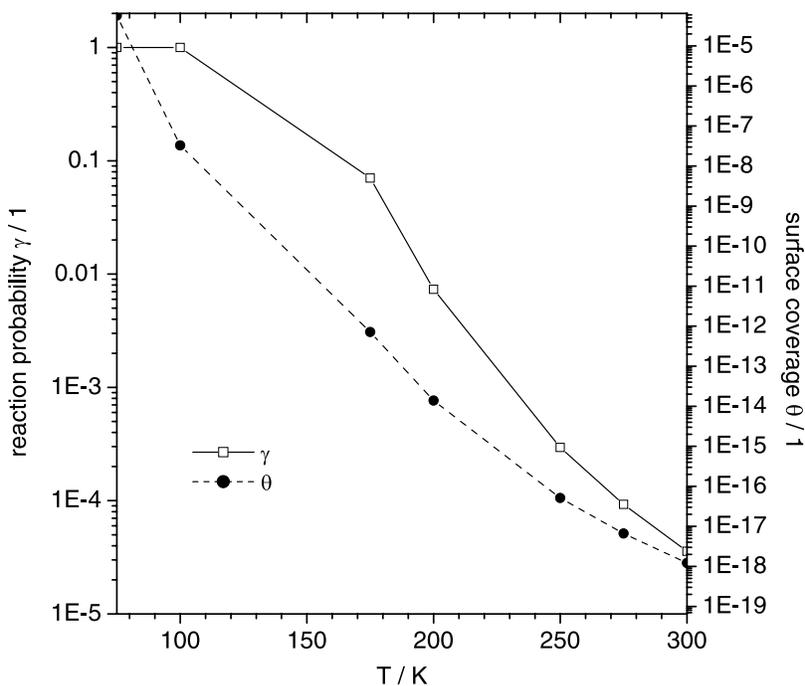


Figure 5. Prediction of the reaction probability γ and surface coverages θ obtained from $k_{\text{ass}}(T)$, the data in Figure 2 and by using equations (7) and (10), respectively, and employing a partial pressure of 10^{-8} Pa ClONO_2 .

Berland *et al.*'s [1997] Figure 11 obtained from a combination of laser-induced thermal desorption and a precursor-mediated adsorption model shows the same S shape as our calculated curve. Even the temperature of 105 K at which γ starts to drop is in agreement. The only apparent discrepancy to this study seems to be the very low trapping coefficient α . From classical trajectory simulations the trapping coefficient α was found on the other hand to be unity [Bolton *et al.*, 1999].

[19] Another experiment has shown no significant temperature dependence between 180 and 200 K, though, where γ was reported to be 0.20 ± 0.05 [Oppliger *et al.*, 1997]. For BrONO₂ hydrolysis a drop of γ from 0.3 (180 K) to 0.025 (210 K) has been found in the same Lausanne laboratory for single-crystal ice surfaces. Similar, but less pronounced negative temperature dependencies for both bulk and condensed ice were also noticed in the same study [Aguzzi, 2001]. The reason why ClONO₂ and BrONO₂ behave differently as a function of temperature is unclear to us and remains subject to further investigation.

[20] In a static secondary ion mass spectrometry study in an ultrahigh vacuum chamber only polarization of chlorine nitrate was observed on ice films held at temperatures as low as 90–120 K but no chemical reaction [Donsig *et al.*, 1999]. According to our rate constant data this interesting finding can be interpreted to arise from the fact that equilibrium was not reached because of the very long equilibration times of the system at these low temperatures. For instance, the rate constant for desorption amounts to less than 10^{-3} s^{-1} at 100 K implying equilibration times of many hours. On the other hand from infrared spectra evidence for the formation of nitric acid, i.e., chemical reaction, was found at temperatures of 140 K [Horn *et al.*, 1998], where the rate constants seen in Figure 2 are of the order of 10^4 s^{-1} , implying equilibration times of seconds.

3.4. Error Estimates

[21] The inherent error for calculating reaction barriers and desorption barriers at G2(MP2) level is estimated to be $1.5 \text{ kcal mol}^{-1}$ [Curtiss *et al.*, 1993]. The reaction barrier increases from $5.61 \text{ kcal mol}^{-1}$ (4:1) to $19.95 \text{ kcal mol}^{-1}$ (3:1) to $26.63 \text{ kcal mol}^{-1}$ (2:1) and to $58.61 \text{ kcal mol}^{-1}$ (1:1) at G2(MP2) level on decreasing the number of water molecules. That is, the uncertainty inherent to the method is much smaller than the difference between different assumed reaction mechanisms. The error in the reaction barrier causes an error in the rate constant because of the proportionality $k \propto \exp(-\Delta G/RT)$. A large fraction of the error in the rate constants cancels out on calculating γ as a ratio of rate constants, especially because of the systematic error in the binding energy inherent to the G2(MP2) calculation. The error in γ is likely to be a few percent, when either $k_{rxn} \gg k_{des}$ or $k_{des} \gg k_{rxn}$. At temperatures where both k_{rxn} and k_{des} are comparable, the errors in k add up in the denominator and result in the worst error, which we presume to be around a factor of 3 around 200 K. That is, the prediction of $\gamma = 0.03$ implies an error range from 0.01 to 0.10. This error is comparable to the experimental uncertainty in the reaction probability determination [Molina *et al.*, 1987; Hanson and Ravishankara, 1991].

[22] Our estimate for γ is slightly lower than the recommended value and the latest determination of γ for chlorine nitrate hydrolysis on ice [Oppliger *et al.*, 1997], but within the

estimated error. It is slightly higher compared to early pioneering studies [Molina *et al.*, 1987; Tolbert *et al.*, 1987; Leu, 1988], which fall in the range 0.007–0.09 at 185–200 K. By using $n = 3$ instead of $n = 4$ for calculating k_{rxn} the estimate is $\gamma \approx 10^{-7}$ at stratospheric conditions, which is clearly lower than all of the experimental results, even those on NAT or sulfuric acid tetrahydrate (SAT) surfaces. Using $n = 1$, thereby calculating the homogeneous reaction, leads to slow reaction and $\gamma = 0$, as also evidenced in the laboratory [Rowland *et al.*, 1986]. By using $n = 5$, $n = 6$ or microsolvated clusters, for which surface reaction becomes barrierless and k_{rxn} much higher, $\gamma = 1$ in the whole temperature range.

4. Conclusions

[23] We have attempted to model the reaction probabilities for chlorine nitrate hydrolysis on an ice surface for the first time a priori without the aid of widely available experimental data. This has been done by performing high-level ab initio calculations on a system composed of one molecule of chlorine nitrate (ClONO₂) and a linear water tetramer (as opposed to the cyclic tetramer [Loerting *et al.*, 1998]). The rate constants of association, surface reaction (including multidimensional tunnelling corrections) and surface desorption have been calculated and used to estimate the heterogeneous reaction probability. Quantum mechanical tunnelling does not play a significant role at tropospheric/stratospheric temperatures in the case of the 4:1 cluster by contrast to the 3:1 and 2:1 clusters. Our prediction for the 4:1 cluster is within the range of the recommended values by IUPAC/JPL, which are based on the available laboratory measurements and chosen so that the effect of the reduction of γ by HNO₃ poisoning is minimized.

[24] The reaction barrier of $5.61 \text{ kcal mol}^{-1}$ obtained at G2(MP2) level of theory agrees with the literature estimate of $3\text{--}7 \text{ kcal mol}^{-1}$ [Tabazadeh *et al.*, 1994]. On increasing the cluster size and/or including microsolvation effects of the ice lattice the reaction barrier disappears [McNamara *et al.*, 1999; Bianco and Hynes, 2006], and on decreasing the cluster size the reaction barrier becomes higher by a factor of 3–10. The former is inconsistent with no detectable chemical reaction on surfaces kept at 90–120 K [Donsig *et al.*, 1999] as well as with a remaining reaction barrier in supersonic molecular beam studies on water clusters containing up to 450 water molecules [Akhmatskaya *et al.*, 1997; Apps *et al.*, 1997]. The latter would imply a negligible reaction probability, which is also inconsistent with experimental data.

[25] Since the experimental observations on an ice surface can be reproduced reasonably with a 4:1 water:chlorine nitrate gas-phase cluster, which shows highly flexible O–O–O angles, we conjecture that for type II PSCs the surface layers are disordered even at 190 K in the presence of ClONO₂. The idea of a liquid-like ice surface at stratospheric temperatures was first postulated in 1994 by Molina to explain HCl adsorption [Molina, 1994]. Evidence for this postulate was provided later by a zeroth-order dependence of the reaction rate constant on HCl partial pressure [Lee *et al.*, 1999] as well as a study of the heterogeneous and photochemistry of ClONO₂ on HCl-doped ice at 181 and 190 K [Faraudo and Weibel, 2001]. Recently it was found that the premelting of ice, i.e., surface disorder, can be strongly enhanced by hydrocarbon contamination [Blum *et al.*,

2002] or induced by HCl traces [McNeill *et al.*, 2006]. The evidence for the migration of acetone into the bulk of an ice film additionally corroborates the possibility of a liquid-like layer under stratospheric conditions in the presence of a variety of impurities [Schaff and Roberts, 1994, 1998]. Further support comes from molecular dynamics simulations showing that the dynamics of the top 2–3 bilayers of water molecules in hexagonal ice at 180–210 K is substantially faster than in the bulk [Bolton and Pettersson, 2000] leading to partial disorder in the uppermost bilayer [Mantz *et al.*, 2000]. This disorder even increases when HCl doses up to the monolayer are deposited on ultrathin (5 bilayers thick) hexagonal ice films supported on MgO(001) as seen from a combined quasi-elastic neutron scattering and molecular dynamics study [Girardet and Toubin, 2001]. Finally, it is hard to imagine how a crystal, with well defined positions of oxygen atoms, would rearrange after a molecule of water is consumed for the reaction. Thus the observations can be reconciled by regarding the ice crystal as being surface-disordered and highly dynamic [Haynes *et al.*, 1992; George and Livingston, 1997]. We want to emphasize that we do not want to imply that a 4:1 mechanism is what happens on polar stratospheric clouds since the microsolvating ice lattice is known to have a strong influence on the reaction barrier, and hence γ [Bianco and Hynes, 2006].

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References

- Abbatt, J. P. D., and M. J. Molina (1992), Heterogeneous interactions of ClONO₂ and HCl on nitric acid trihydrate at 202 K, *J. Phys. Chem.*, *96*, 7674–7679.
- Adamson, A. W. (1997), *Physical Chemistry of Surfaces*, 6th ed., John Wiley, Hoboken, N. J.
- Aguzzi, A. (2001), Atmospheric heterogeneous reactions of chlorine and bromine containing molecules: A laboratory study, Ph.D. thesis, Ecole Polytech. Féd. de Lausanne, Lausanne, Switzerland. (Available at <http://lpas.epfl.ch/>)
- Akhmatskaya, E. V., C. J. Apps, I. H. Hillier, A. J. Masters, I. J. Palmer, N. E. Watt, M. A. Vincent, and J. C. Whitehead (1997), Hydrolysis of SO₃ and ClONO₂ in water clusters: A combined experimental and theoretical study, *J. Chem. Soc., Faraday Trans.*, *93*, 2775–2779.
- Apps, C. J., N. E. Watt, and J. C. Whitehead (1997), Adsorption of some atmospherically important molecules onto large water clusters: SO₂, SO₃, HCl, and ClONO₂, *Isr. J. Chem.*, *37*, 419–425.
- Atkinson, R., E. C. Tuazon, H. M. Leod, S. M. Aschmann, and A. M. Winer (1986), The gas-phase reaction of chlorine nitrate with water vapor, *Geophys. Res. Lett.*, *13*, 117–120.
- Becke, A. D. (1993), Density-functional thermochemistry. III. The role of exact exchange, *J. Chem. Phys.*, *98*, 5648–5652.
- Becke, A. D. (1997), Density-functional thermochemistry. V. Systematic optimization of exchange-correlation functionals, *J. Chem. Phys.*, *107*, 8554–8560.
- Berland, B. S., M. A. Tolbert, and S. M. George (1997), Surface sensitive studies of the reactive uptake of chlorine nitrate on ice, *J. Phys. Chem. A*, *101*, 9954–9963.
- Bianco, R., and J. T. Hynes (1998), Ab initio model study of the mechanism of chlorine nitrate hydrolysis on ice, *J. Phys. Chem. A*, *102*, 309–314.
- Bianco, R., and J. T. Hynes (2006), Heterogeneous reactions important in atmospheric ozone depletion: A theoretical perspective, *Accounts Chem. Res.*, in print.
- Bianco, R., B. J. Gertner, and J. T. Hynes (1998), Proton transfer reactions at the surface of ice, heterogeneous reactions involved in stratospheric ozone depletion, *Ber. Bunsen Ges. Phys. Chem.*, *102*, 518–526.
- Bianco, R., W. H. Thompson, A. Morita, and J. T. Hynes (2001), Is the H₂OCl⁺ ion a viable intermediate for the hydrolysis of ClONO₂ on ice surfaces?, *J. Phys. Chem. A*, *105*, 3132–3139.
- Billing, G. D. (2000), *Dynamics of Molecule Surface Interactions*, John Wiley, Hoboken, N. J.
- Bluhm, H., D. F. Ogletree, C. S. Fadley, Z. Hussain, and M. Salmeron (2002), The premelting of ice studied with photoelectron spectroscopy, *J. Phys. Condens. Mat.*, *14*, L227–L233.
- Bolton, K., and J. B. C. Pettersson (2000), A molecular dynamics study of the long time Ih surface dynamics, *J. Phys. Chem. B*, *104*, 1590–1595.
- Bolton, K., M. Swanberg, and J. B. C. Pettersson (1999), Classical trajectory study of Ar-ice collision dynamics, *J. Chem. Phys.*, *110*, 5380–5391.
- Carlsaw, K. S., and T. Peter (1997), Uncertainties in reactive uptake coefficients for solid stratospheric particles: 1. Surface chemistry, *Geophys. Res. Lett.*, *24*, 1743–1746.
- Carlsaw, K. S., T. Peter, and R. Mueller (1997), Uncertainties in reactive uptake coefficients for solid stratospheric particles: 2. Effect on ozone depletion, *Geophys. Res. Lett.*, *24*, 1747–1750.
- Chu, L. T., M.-T. Leu, and L. F. Keyser (1993), Heterogeneous reactions of HOCl + HCl → Cl₂ + H₂O and ClONO₂ + HCl → Cl₂ + HNO₃ on ice surfaces at polar stratospheric conditions, *J. Phys. Chem.*, *97*, 12,798–12,804.
- Chuang, Y.-Y., *et al.* (1999), Polyrate8.6, Univ. of Minn., Minneapolis.
- Corchado, J. C., E. L. Coiti no, Y.-Y. Chuang, and D. G. Truhlar (1999), Gaussrate8.6, University of Minn., Minneapolis.
- Curtiss, L. A., K. Raghavachari, and J. A. Pople (1993), Gaussian-2 theory using reduced Møller-Plesset orders, *J. Chem. Phys.*, *98*, 1293–1298.
- Del Negro, L. A., *et al.* (1997), Evaluating the role of NAT, NAD, and liquid H₂SO₄/H₂O/HNO₃ solutions in antarctic polar stratospheric cloud aerosol: Observations and implications, *J. Geophys. Res.*, *102*, 13,255–13,282.
- DeMore, W. B., S. P. Sander, D. M. Golden, R. F. Hampson, M. J. Kurylo, C. J. Howard, A. R. Ravishankara, C. E. Kolb, and M. J. Molina (1997), Chemical kinetics and photochemical data for use in stratospheric modeling, *JPL Publ.*, *97-4*.
- Donsig, H. A., D. Herridge, and J. C. Vickerman (1999), Static SIMS studies of reactions on mimics of polar stratospheric clouds III: Mechanism of chlorine nitrate decomposition and reaction, *J. Phys. Chem. A*, *103*, 9211–9220.
- Faraudo, G., and D. E. Weibel (2001), Effect of HCl addition and temperature on the heterogeneous chemistry and photochemistry of ClONO₂ adsorbed on ice crystals, *Prog. React. Kinet. Mech.*, *26*, 179–199.
- Finlayson-Pitts, B. J., M. J. Ezell, and J. N. Pitts Jr. (1989), Formation of chemically active chlorine compounds by reactions of atmospheric sodium chloride particles with gaseous dinitrogen pentoxide and chlorine nitrate, *Nature*, *337*, 241–244.
- Frisch, M. J., *et al.* (1998), Gaussian 98, Revision A.7, Gaussian, Inc., Pittsburgh, Pa.
- Geiger, F. M., C. D. Pibel, and J. M. Hicks (2001), The hydrolysis of chlorine nitrate on ice is autocatalytic, *J. Phys. Chem. A*, *105*, 4940–4945.
- George, S. M., and F. M. Livingston (1997), Dynamic ice surface in the polar stratosphere, *Surface Rev. Lett.*, *4*, 771–780.
- Gilligan, J. J., and A. W. Castleman (2001), Direct experimental evidence for reactions between dissolved acid halide and chlorine nitrate, *J. Phys. Chem. A*, *105*, 1028–1032.
- Gilligan, J. J., D. J. Moody, and A. W. Castleman (2000), Reactions of protonated water clusters with chlorine nitrate revisited, *Z. Phys. Chem.*, *214*, 1383–1391.
- Girardet, C., and C. Toubin (2001), Molecular atmospheric pollutant adsorption on ice: A theoretical survey, *Surface Sci. Rep.*, *44*, 159–238.
- Grassian, V. H. (2001), Heterogeneous uptake and reaction of nitrogen oxides and volatile organic compounds on the surface of atmospheric particles including oxides, carbonates, soot and mineral dust: Implications for the chemical balance of the troposphere, *Int. Rev. Phys. Chem.*, *20*, 467–548.
- Hammond, G. S. (1955), A correlation of reaction rates, *J. Am. Chem. Soc.*, *77*, 334–338.
- Hanson, D. R., and A. R. Ravishankara (1991), The reaction probabilities of ClONO₂ and N₂O₅ on polar stratospheric cloud materials, *J. Geophys. Res.*, *96*, 5081–5090.
- Hanson, D. R., and A. R. Ravishankara (1992), Investigation of the reactive and nonreactive processes involving ClONO₂ and HCl on water and nitric acid doped ice, *J. Phys. Chem.*, *96*, 2682–2691.
- Hanson, D. R., and A. R. Ravishankara (1993), Response to "Comment on porosities of ice films used to simulate stratospheric cloud surfaces", *J. Phys. Chem.*, *97*, 2802–2803.
- Haynes, D. R., N. J. Tro, and S. M. George (1992), Condensation and evaporation of water on ice surfaces, *J. Phys. Chem.*, *96*, 8502–8509.
- Hobbs, P. V. (1974), *Ice Physics*, Clarendon, Oxford, U. K.
- Horn, A. B., J. R. Sodeau, T. B. Roddis, and N. A. Williams (1998), Low temperature reaction of chlorine nitrate with water ice: Formation of molecular nitric acid, *J. Chem. Soc. Farad. Trans.*, *94*, 1721–1724.

- Keyser, L. F., S. B. Moore, and M.-T. Leu (1991), Surface reaction and pore diffusion in flow-tube reactors, *J. Phys. Chem.*, *95*, 5496–5502.
- Koch, T. G., S. F. Banham, J. R. Sodeau, A. B. Horn, M. R. S. McCoustra, and M. A. Chesters (1997), Mechanisms for the heterogeneous hydrolysis of hydrogen chloride, chlorine nitrate and dinitrogen pentoxide on water-rich atmospheric particle surfaces, *J. Geophys. Res.*, *102*, 1513–1522.
- Kolb, C. E., D. R. Worsnop, M. S. Zahniser, P. Davidovits, L. F. Keyser, M.-T. Leu, M. J. Molina, D. R. Hanson, and A. R. Ravishankara (1995), Laboratory Studies of Atmospheric Heterogeneous Chemistry, in *Progress and Problems in Atmospheric Chemistry, Adv. Ser. Phys. Chem.*, vol. 3, pp. 771–875, World Scientific, Hackensack, N. J.
- Langmuir, I. (1922), Chemical reaction on surfaces, *Trans. Faraday Soc.*, *17*, 607–620.
- Lee, S.-H., D. C. Leard, R. Zhang, L. T. Molina, and M. J. Molina (1999), The HCl+ClONO₂ reaction on various water ice surfaces, *Chem. Phys. Lett.*, *315*, 7–11.
- Leu, M.-T. (1988), Laboratory studies of sticking coefficients and heterogeneous reactions important in the antarctic stratosphere, *Geophys. Res. Lett.*, *15*, 17–20.
- Leu, M. T., S. B. Moore, and L. F. Keyser (1991), Heterogeneous reactions of chlorine nitrate and hydrogen chloride on type I polar stratospheric clouds, *J. Phys. Chem.*, *95*, 7763–7771.
- Loerting, T., and K. R. Liedl (2001), The reaction rate constant of chlorine nitrate hydrolysis, *Chem. Eur. J.*, *7*, 1662–1669.
- Loerting, T., K. R. Liedl, and B. M. Rode (1998), Predictions of rate constants and estimates for tunneling splittings of concerted proton transfer in small cyclic water clusters, *J. Chem. Phys.*, *109*, 2672–2679.
- Mantz, Y. A., F. M. Geiger, L. T. Molina, M. J. Molina, and B. L. Trout (2000), First-principles molecular-dynamics study of surface disordering of the (0001) face of hexagonal ice, *J. Chem. Phys.*, *113*, 10,733–10,743.
- McNamara, J. P., and I. H. Hillier (1999), Exploration of the mechanism of the hydrolysis of chlorine nitrate in small water clusters using electronic structure methods, *J. Phys. Chem. A*, *103*, 7310–7321.
- McNamara, J. P., and I. H. Hillier (2000), The facile hydrolysis of chlorine nitrate in aqueous sulfate aerosols, *Chem. Phys. Lett.*, *328*, 492–499.
- McNamara, J. P., and I. H. Hillier (2001), Mechanism of the hydrolysis of halogen nitrates in small water clusters studied by electronic structure methods, *J. Phys. Chem. A*, *105*, 7011–7024.
- McNamara, J. P., G. Tresadern, and I. H. Hillier (1999), The facile decomposition of chlorine nitrate in small water clusters, *Chem. Phys. Lett.*, *310*, 265–270.
- McNeill, V. F., T. Loerting, F. M. Geiger, B. L. Trout, and M. J. Molina (2006), Hydrogen chloride induced surface disordering on ice, *Proc. Natl. Acad. Sci. U.S.A.*, *103*, 9422–9427.
- Molina, M. J. (1994), The probable role of stratospheric ‘ice’ clouds: Heterogeneous chemistry of the ‘ozone hole’, in *The Chemistry of the Atmosphere: Its Impact on Global Change*, pp. 27–38, Blackwell Sci., Malden, Mass.
- Molina, M. J., T.-L. Tso, L. T. Molina, and F. C.-Y. Wang (1987), Antarctic stratospheric chemistry of chlorine nitrate, hydrogen chloride, and ice: Release of active chlorine, *Science*, *238*, 1253–1257.
- Molina, M. J., R. Zhang, P. J. Wooldridge, J. R. McMahon, J. E. Kim, H. Y. Chang, and K. D. Beyer (1993), Physical chemistry of the H₂SO₄/HNO₃/H₂O system: Implications for polar stratospheric clouds, *Science*, *261*, 1418–1423.
- Molina, M. J., L. T. Molina, and C. E. Kolb (1996), Gas-phase and heterogeneous chemical kinetics of the troposphere and stratosphere, *Annu. Rev. Phys. Chem.*, *47*, 327–367.
- Moore, S. B., L. F. Keyser, M.-T. Leu, R. P. Turco, and R. H. Smith (1990), Heterogeneous reactions on nitric acid trihydrate, *Nature*, *345*, 333–335.
- Mozurkewich, M. (1993), Effect of competitive adsorption on polar stratospheric cloud reactions, *Geophys. Res. Lett.*, *20*, 355–358.
- Oppliger, R., A. Allan, and M. J. Rossi (1997), Real-time kinetics of the uptake of ClONO₂ on ice and in the presence of HCl in the temperature range 160 K ≤ T ≤ 200 K, *J. Phys. Chem. A*, *101*, 1903–1911.
- Pilling, M. J. (1995), *Reaction Kinetics*, Oxford Univ. Press, New York.
- Robinson, G. N., D. R. Worsnop, J. T. Jayne, C. E. Kolb, and P. Davidovits (1997), Heterogeneous uptake of ClONO₂ and N₂O₅ by sulfuric acid solutions, *J. Geophys. Res.*, *102*, 3583–3601.
- Rowland, F. S., H. Sato, H. Khwaja, and S. M. Elliott (1986), The hydrolysis of chlorine nitrate and its possible atmospheric significance, *J. Phys. Chem.*, *90*, 1985–1988.
- Sander, S. P., et al. (2000), Chemical kinetics and photochemical data for use in stratospheric modeling: Update of key reactions, *JPL Publ.*, 00-3.
- Schaff, J. E., and J. T. Roberts (1994), Structure sensitivity in the surface chemistry of ice: Acetone adsorption on amorphous and crystalline ice films, *J. Phys. Chem.*, *98*, 6900–6902.
- Schaff, J. E., and J. T. Roberts (1998), The adsorption of acetone on films of amorphous and crystalline ice, *Langmuir*, *14*, 1478–1486.
- Schindler, T., C. Berg, G. Niedner-Schatteburg, and V. E. Bondyney (1996), Heterogeneously catalyzed hydrolysis of chlorine nitrate: Fourier-transform ion cyclotron resonance investigations of stratospheric chemistry, *J. Chem. Phys.*, *104*, 3998–4004.
- Schindler, T., C. Berg, K. Fritz, U. Achatz, M. Beyer, G. Niedner-Schatteburg, and V. E. Bondyney (2000), A case study on molecular clusters and atmospheric aerosols: Chlorine nitrate hydrolysis, in *Recent Theoretical and Experimental Advances, in Hydrogen Bonded Clusters, NATO ASI Ser. C*, vol. 561, pp. 389–391, Springer, New York.
- Shi, Q., J. T. Jayne, C. E. Kolb, D. R. Worsnop, and P. Davidovits (2001), Kinetic model for reaction of ClONO₂ with H₂O and HCl and HOCl with HCl in sulfuric acid solutions, *J. Geophys. Res.*, *106*, 4,259–4,274.
- Solomon, S. (1990), Progress towards a quantitative understanding of antarctic ozone depletion, *Nature*, *347*, 347–354.
- Solomon, S., R. R. Garcia, F. S. Rowland, and D. J. Wuebbles (1986), On the depletion of antarctic ozone, *Nature*, *321*, 755–758.
- Solomon, S., S. Borrmann, R. R. Garcia, R. Portmann, L. Thomason, L. R. Poole, D. Winker, and M. P. McCormick (1997), Heterogeneous chlorine chemistry in the tropopause region, *J. Geophys. Res.*, *102*, 21,411–21,429.
- Tabazadeh, A., and R. P. Turco (1993), A model for heterogeneous chemical processes on the surfaces of ice and nitric acid trihydrate particles, *J. Geophys. Res.*, *98*, 12,727–12,740.
- Tabazadeh, A., R. P. Turco, K. Drdla, M. Z. Jacobson, and O. B. Toon (1994), A study of type I polar stratospheric cloud formation, *Geophys. Res. Lett.*, *21*, 1619–1622.
- Tolbert, M. A., and O. B. Toon (2001), Solving the PSC mystery, *Science*, *292*, 61–63.
- Tolbert, M. A., M. J. Rossi, R. Malhotra, and D. M. Golden (1987), Reaction of chlorine nitrate with hydrogen chloride and water at antarctic stratospheric temperatures, *Science*, *238*, 1258–1260.
- Tolbert, M. A., M. J. Rossi, and D. M. Golden (1988), Heterogeneous interactions of chlorine nitrate, hydrogen chloride, and nitric acid with sulfuric acid surfaces at stratospheric temperatures, *Geophys. Res. Lett.*, *15*, 847–850.
- Toon, O. B., A. Tabazadeh, E. V. Browell, and J. Jordan (2000), Analysis of lidar observations of Arctic polar stratospheric clouds during January 1989, *J. Geophys. Res.*, *105*, 20,589–20,616.
- Truhlar, D. G. (1995), Direct dynamics method for the calculation of reaction rates, in *The Reaction Path in Chemistry: Current Approaches and Perspectives*, edited by D. Heidrich, pp. 229–255, Springer, New York.
- Truhlar, D. G., and B. C. Garrett (1984), Variational transition state theory, *Annu. Rev. Phys. Chem.*, *35*, 159–189.
- Truhlar, D. G., and M. S. Gordon (1990), From force fields to dynamics: Classical and quantal paths, *Science*, *249*, 491–498.
- Truhlar, D. G., A. D. Isaacson, and B. C. Garrett (1985), Generalized transition state theory, in *Theory of Chemical Reaction Dynamics*, edited by M. Baer, pp. 65–137, CRC Press, Boca Raton, Fla.
- Xu, S. C., and X. S. Zhao (1999), Theoretical investigation of the reaction of ClONO₂ with H₂O on water clusters, *J. Phys. Chem. A*, *103*, 2100–2106.
- Ying, L., and X. Zhao (1997), Theoretical studies of XONO₂-H₂O (X = Cl, H) complexes, *J. Phys. Chem. A*, *101*, 6807–6812.
- Zhang, R., P. J. Wooldridge, and M. J. Molina (1993), Vapor pressure measurements for sulfuric acid/nitric acid/water and sulfuric acid/hydrochloric acid/water systems: Incorporation of stratospheric acids into background sulfate aerosols, *J. Phys. Chem.*, *97*, 8541–8548.
- Zhang, R., J. T. Jayne, and M. J. Molina (1994), Heterogeneous interactions of nitryl hypochlorite and hydrogen chloride with sulfuric acid tetrahydrate: Implications for the stratosphere, *J. Phys. Chem.*, *98*, 867–874.
- Zhang, R., M.-T. Leu, and L. F. Keyser (1995), Hydrolysis of N₂O₅ and ClONO₂ on the H₂SO₄/HNO₃/H₂O ternary solutions under stratospheric conditions, *Geophys. Res. Lett.*, *22*, 1493–1496.
- Zondlo, M. A., S. B. Barone, and M. A. Tolbert (1998), Condensed-phase products in heterogeneous reactions: N₂O₅, ClONO₂, and HNO₃ reacting on ice films at 185 K, *J. Phys. Chem. A*, *102*, 5735–5748.
- Zondlo, M. A., P. K. Hudson, A. J. Prenni, and M. A. Tolbert (2000), Chemistry and microphysics of polar stratospheric clouds and cirrus clouds, *Annu. Rev. Phys. Chem.*, *51*, 473–499.

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